

# Strong Earthquake Sequences in Greece during 2008-2014: Moment Tensor Inversions and Fault Plane Discrimination

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**How to cite this paper:** Moshou, A. (2020) Strong Earthquake Sequences in Greece during 2008-2014: Moment Tensor Inversions and Fault Plane Discrimination. *Open Journal of Earthquake Research*, 9, 323-348. <https://doi.org/10.4236/ojer.2020.94019>

**Received:** May 1, 2020

**Accepted:** July 31, 2020

**Published:** August 3, 2020

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## Abstract

As is well known, Greece has a significant number of earthquakes each year. In recent years, several earthquakes have occurred in Greece. For this scope, a methodology was used to determine the source parameters. This methodology is based on minimizing the difference between the observed and the synthetic waveforms, using the method Source Parameters Calculation—SPCa [1]. The source parameters, using the proposed methodology, are calculated by comparing observed seismograms and synthetic by inverting data. The synthetics are calculated using the reflectivity method (Kennett, 1983) as implemented by Randall *et al.* (1994) for a given earth structure. This study includes inversion results for the strongest events that occurred in Greece from 2008 to 2014. For the same events calculated the main fault plane, using the method of Hypocenter Centroid-plot (HC-plot) [2] [3]. This methodology is a simple geometrical method based on the combination between the hypocentral position and the two possible fault planes.

## Keywords

Synthetic Seismograms, Moment Tensor Inversion, Focal Mechanism, Regional Data, Nodal Planes

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## 1. Introduction

The seismic moment tensor constitutes the most important source parameter since it describes in a first-order approximation the equivalent forces applied on a fault plane and can be calculated by body wave modeling. The moment tensor as a mathematical description of equivalent forces and moments is used to study the source processes. The propagation and the source effects characterize varia-

tions of the observed seismograms. Mathematically, each one of these effects can be calculated to generate synthetic seismograms that can be compared directly to the corresponding observed ones. The best solution is obtained by the minimization of the difference between the observed and the synthetic seismograms (Aki and Richards, 1980).

General a seismic source can be representing by a symmetric square matrix ( $3 \times 3$ ) with 6 independent elements. Eigenvalue and eigenvector analysis of the moment tensor can be used to determine the components of the moment tensor. All the mathematical expressions described analytically in the studies [4] [5].

The linear dipole is compensated by two other linear dipoles along with the other two perpendicular directions. The eigenvalues describe the isotropic component of the moment tensor. In the case that the sum of the eigenvalues is vanished the applied forces constitute a pure double couple source. In this case, the seismic moment tensor has only deviatoric components [5] [6]. In general, a complete moment tensor is the superposition of the two vector dipoles (DC and CLVD) and isotropic component [5] [6]. In the case of an earthquake, the isotropic part is zero.

## 2. Applied Methodology and Preparation of Data

Seismological digital broadband data from the Hellenic Unified Seismological Network (HUSN) were collected and analyzed to calculate the source parameters of the strongest earthquakes that occurred in the Greece area, for the last six years. For this purpose, a methodology based on a moment tensor inversion was used, using the software of Ammon [7]. This method calculates synthetic seismograms directly compared with the observed ones for a given velocity structure. The reflectivity method of Kennett [8] as implemented by Randall [9] was applied to determine the Green Functions, initially calculated for different depths by the analyst. Iterative inversions were performed at depth intervals of 5 km followed by a finer one of 1 - 2 km steps around the depth exhibiting the lowest misfit.

Regional data at least five broadband stations, at different azimuthal coverage and epicentral distances less than  $3^\circ$ , equipped with three components seismometers, were selected and analyzed. The preparation of the data, includes the deconvolution of instrument response, following the velocity was integrated to displacement and finally, the horizontal components rotated to radial and transverse ones. Then the method uses the long period part of the signal to invert.

A bandpass filter is applied both on the observed waveforms and synthetics, having a fixed length of 70 sec. The inversion results indicate that inverting waveforms longer than 70 sec resulted in higher misfits. The quality of the results of moment tensor solutions can be evaluated by considering the average misfit and the compensated linear vector dipole (CLVD). For each solution, there is a quality code that consists of the letters A - D, for the minimum misfit

and between the numbers 1 - 4 for the percent of CLVD [9].

Finally, for all the previous events we determined the main fault plane that activated, using the Hypocenter Centroid-plot (HC-plot). In this methodology is putting hypocenter and centroid of the 3 components local waveform inversion is located in three-dimensional space, and later calculate its distance to both faulting planes. If the hypocenter is located on one of the two plane faults, so the fault plane is the real fault plane. If the hypocenter is not located in one of the two-fault planes the real fault plane is the closest one to the hypo-center [2] [3].

### 3. Applications

The proposed methodology was applied to the largest earthquakes that occurred recently in Greece. The first two applications concern the 2006 Kythira and the 2008 Leonidio earthquakes (deep events), while the last five concern the seismic sequences that occurred in Greece last year.

On February 2008, three strong events ( $M_w = 6.7, 6.1$  and  $6.0$ ) occurred South of Methoni, at a segment of the Hellenic arc which was not activated during the instrumental period. This sequence was followed by a large number of aftershocks, the strongest of which were processed to calculate their source parameters. On the other hand, the aftershock distribution of the 2008 Andravida ( $M_w = 6.4$ ) earthquake extended to an area significantly larger than the one expected according to the magnitude of the main event. Also, the source parameters for the three seismic sequences in Limnos Island (08/01/2013), in Kallidromon Mountain (Central Greece, 07/08/2013) and finally the most recent in Kefallinia Island (26/01/2014) were calculated. The obtained source parameters were compared to the seismotectonic characteristics of each seismogenic area. These events occurred in different seismotectonic settings, the fact that permits us to evaluate the reliability of the method.

#### 3.1. The $M_w = 6.6$ , 2006 Kythira Earthquake

On 8 January 2006 (11:34 GMT) an earthquake of Magnitude  $M_w = 6.6$  occurred close to the Northeast coast of Kythira Island (Southern Greece) causing some damages landslides and rockfalls. The hypocenter as estimated by the National Observatory of Athens,

[http://bbnet.gein.noa.gr/alerts\\_manual/2006/01/evman060108113454\\_info.html](http://bbnet.gein.noa.gr/alerts_manual/2006/01/evman060108113454_info.html)

was calculated  $36.214^\circ\text{N}$ ,  $23.406^\circ\text{E}$ , and the focal depth at 69 Km. This event is located in the Hellenic subduction zone which characterizes the Southern part of Greece. The shock was felt in a spatially extended area that covered Greece, Italy, Turkey, Egypt, Cyprus, Israel, Syria, Jordan, and Lebanon. This event is one of the largest earthquakes, after the event of 1903 ( $M = 7.3$ ), which occurred in the same area. Previous works indicate the existence of a seismic gap in this region [10]. The major part of the seismic activity in this region is related to the active subduction zone along the Hellenic Arc, as well as the backarc area. Consequently, the area presents complex deformation, where the southern part of the

Aegean region is moving towards the southwest at approximately 40 mm/yr [11] [12]. The area is characterized by active normal faults (Danamos, 1992) with an almost vertical orientation concerning the subduction zone. The main event was followed by a small number of aftershocks, as it appeared in **Figure 1**.

Thrust type faulting was revealed after applying moment tensor inversion. The obtained focal mechanism is strike = 205°, dip = 48° and rake = 59° with a seismic moment equal to  $8.4 \times 10^{25}$  dyn-cm, and a focal depth equal to 65 km (**Figure 2**).

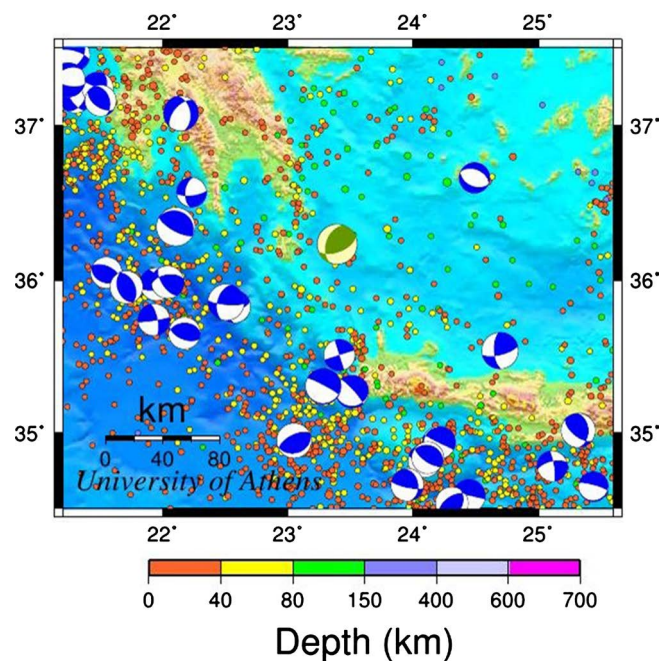
The obtained focal mechanism calculated using data in regional distances is in agreement with the one proposed by the Harvard CMT solution, using teleseismic data.

Using manually locations of HRV, NOA and UOA and their respectively Moment Tensors (**Table 1**) with the varied depth we obtain H-C consistent solution.

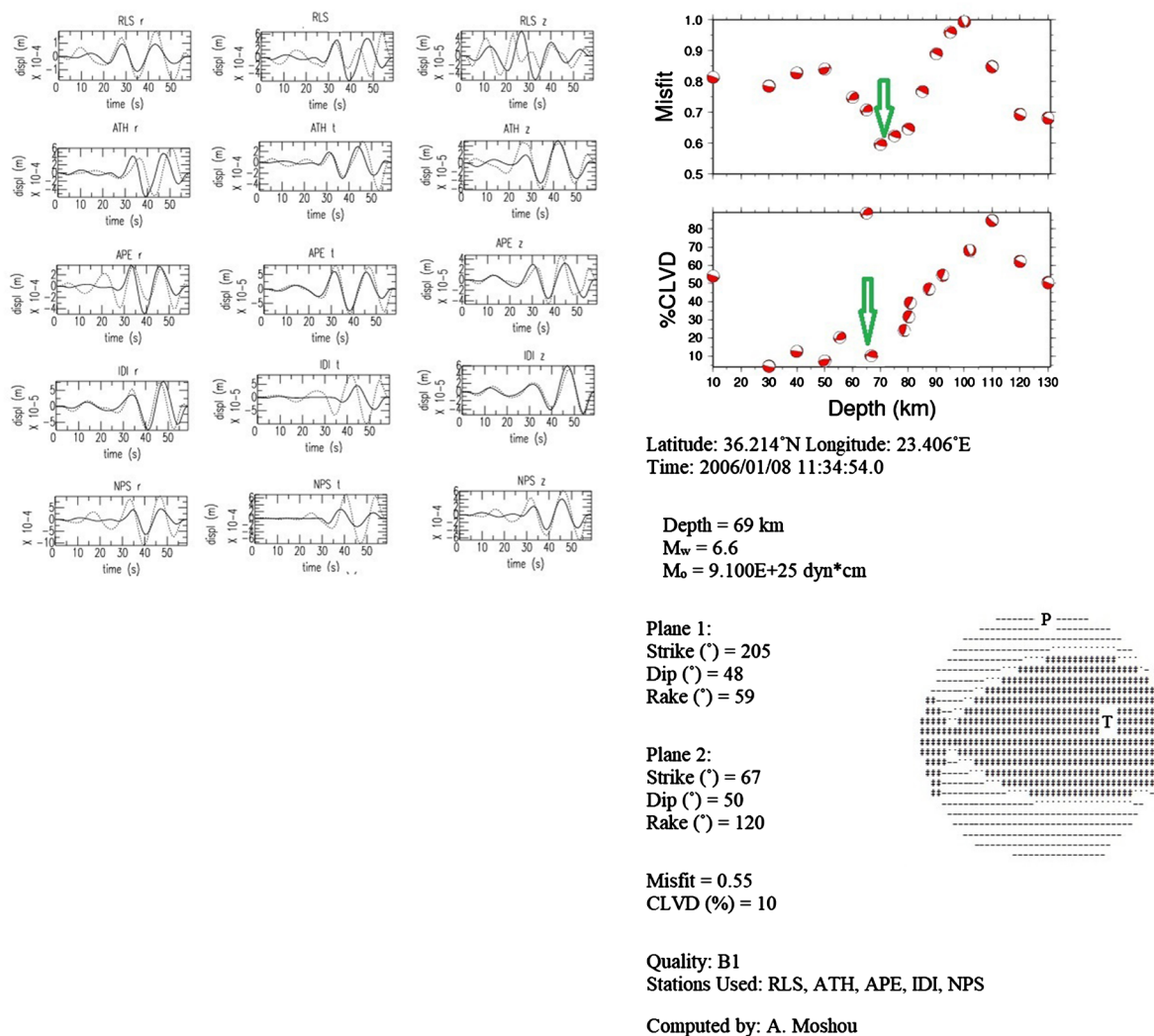
The fault plane is the nodal plane with strike = 205°, dip = 48° and strike = 59°. The distance of the hypocenter from this plane is 7.5 km, while the distance from the other plane is 22.30 km. The distance between Hypocenter and Centroid is 33 km (**Figure 3**).

### 3.2. The Mw = 6.0, 2008 Leonidio Earthquake

On January 6, 2008 (05:14, UTC) a strong earthquake happened near the Leonidio



**Figure 1.** Hypocenter of aftershocks detected and localized using data from Geofon network, for the time interval 1990-2006 is shown by oranges, yellow and green circles respectively to the depth scale. The focal mechanism of the main event is represented by a green beach ball while blue color appears all the other focal mechanisms in the same region, [http://www.geophysics.geol.uoa.gr/imageseis/eqs/2006/20060108\\_kyth/earthquake\\_large\\_en.html#meca](http://www.geophysics.geol.uoa.gr/imageseis/eqs/2006/20060108_kyth/earthquake_large_en.html#meca).



**Figure 2.** Moment tensor solution of the 8 January 2006 (11:34 GMT) earthquake. The selected solution is highlighted with the green arrow in the misfit/CLVD-versus-depth diagrams. Down of this are appeared the summary of the solution and the corresponding beach ball. To the left of misfit/CLVD diagrams observed and synthetic displacement waveforms (continuous and dotted lines respectively) are shown, at the inverted stations for the radial, tangential and vertical components. At the lower part of the figure the summary of the solution and the fault plane solution as lower hemisphere equal-area projection, are depicted.

**Table 1.** Manual Locations and focal mechanisms by various agencies, Sources: HARV, UOA solution.

Agency	Lat (°)	Lon (°)	Depth (km)	Mw	Strike 1 (°)	Dip 1 (°)	Rake 1 (°)	Strike 2 (°)	Dip 2 (°)	Rake 2 (°)
This study	36.21	23.40	69	6.6	205	48	59	67	50	120
HARV	35.93	23.29	64	6.7	201	44	55	66	55	119
UOA	36.23	23.39	65	6.6	210	36	50	76	63	115

town, Southern Peloponnesus. The geographical coordinates, as they calculated by National Observatory of Athens found, 37.114°N, 22.775°E and the depth at 86 km. There were no injuries or damages reported. The main event was followed by a small number of aftershocks since it is an intermediate-depth earthquake (Figure 4).

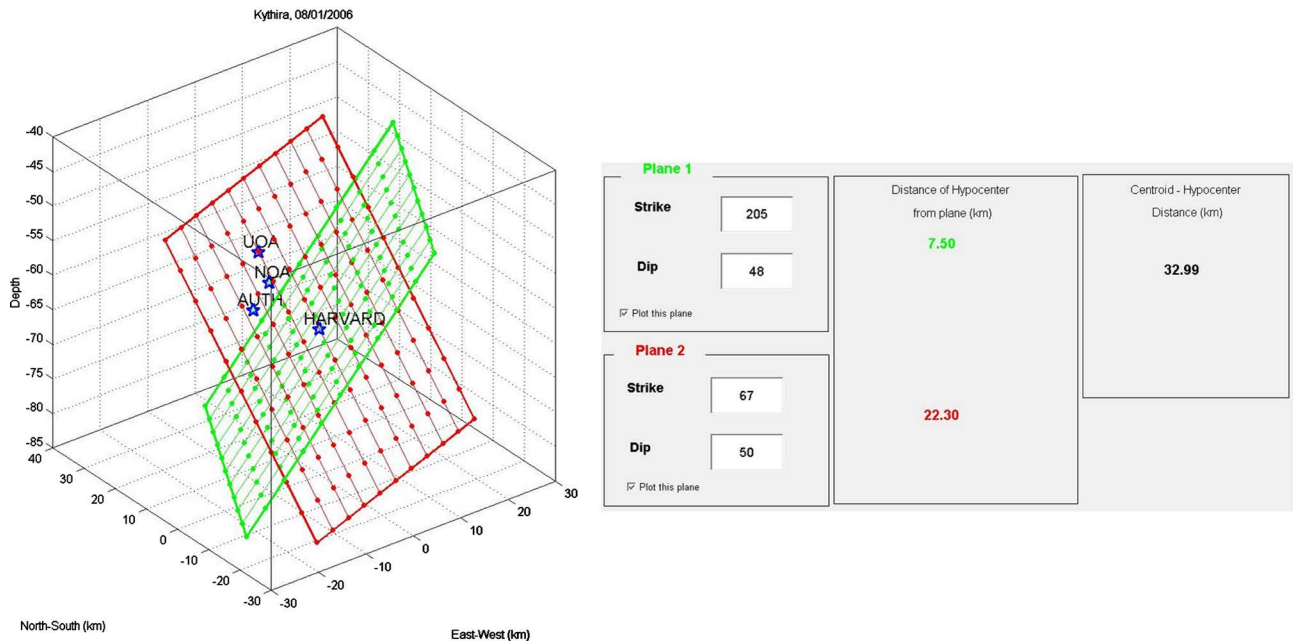


Figure 3. H-C Plot using the NOA and UOA epicenters (with varying depth), and combining it with the CMT solution of HRV.

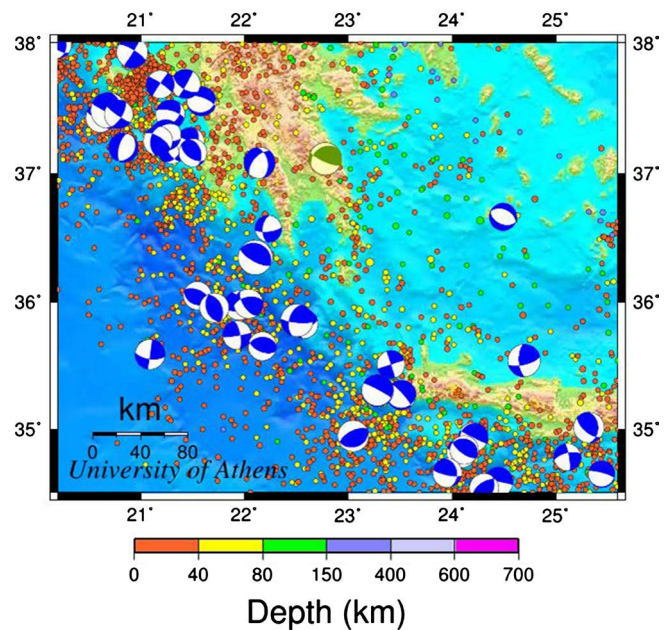


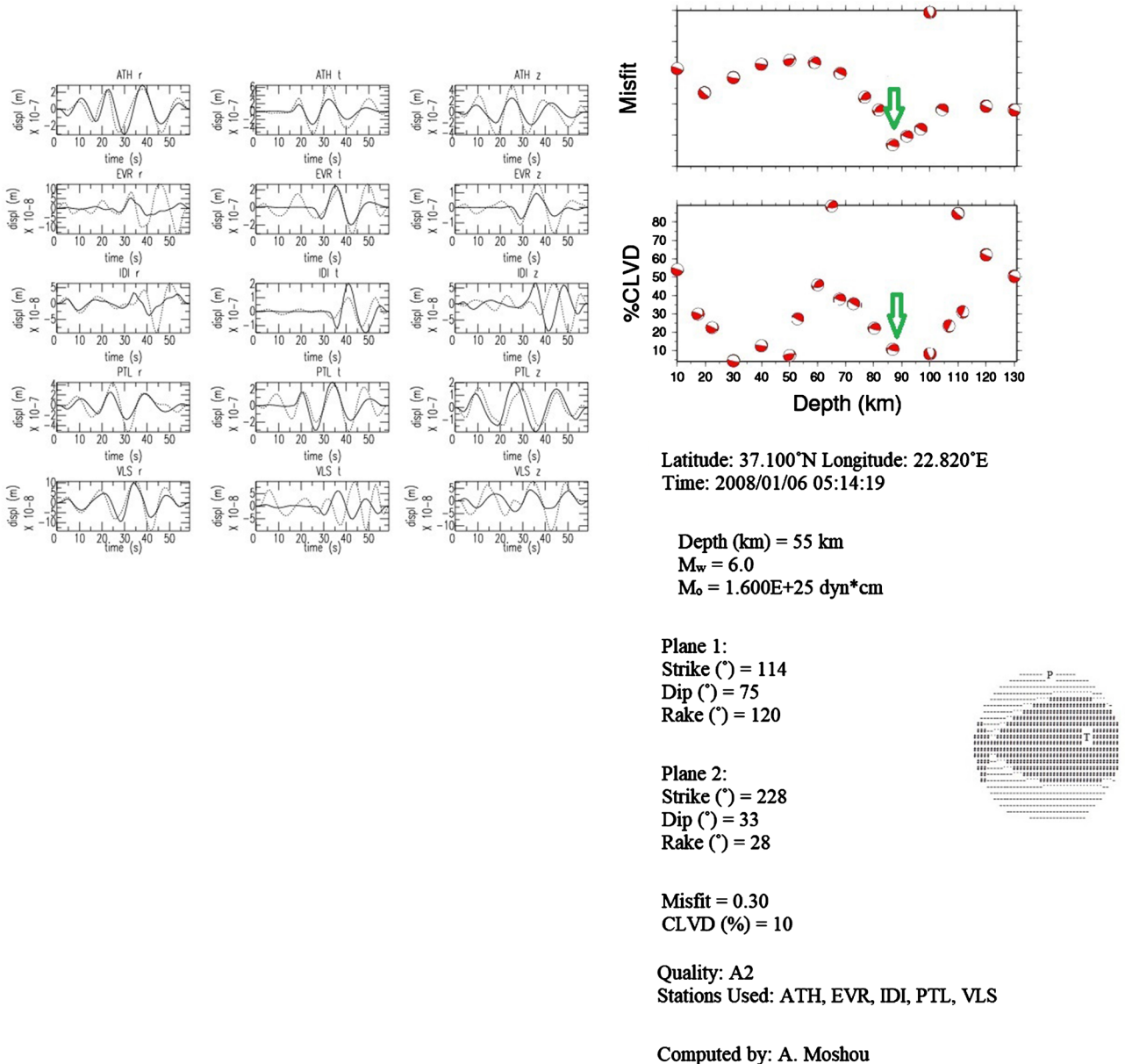
Figure 4. Hypocenter of aftershocks detected and localized using broadband data from Orfeus Institute and Geofon networks, is shown by oranges, yellow and green circles respectively to the depth scale. The focal mechanism of the main event is represented by a green beach ball while blue color appears all the other focal mechanisms in the same region, [http://www.geophysics.geol.uoa.gr/imageseis/eqs/2008/20080106\\_leo/earthquake\\_large\\_e\\_n.html#meca](http://www.geophysics.geol.uoa.gr/imageseis/eqs/2008/20080106_leo/earthquake_large_e_n.html#meca).

Three-component seismological data from stations belonging to the ORFEUS Institute were used. The inversion procedure provided a thrust type faulting with source parameters: strike = 114°, dip = 75° and rake = 120°. The depth is calculated at 85 km and the seismic moment  $M_0 = 1.6 \times 10^{25}$  dyn-cm. The obtained

result showed a good fit between the observed and the synthetic waveforms (Figure 5).

To apply the HC-plot method, the following 6 hypocenters manually locations (Table 2) were considered.

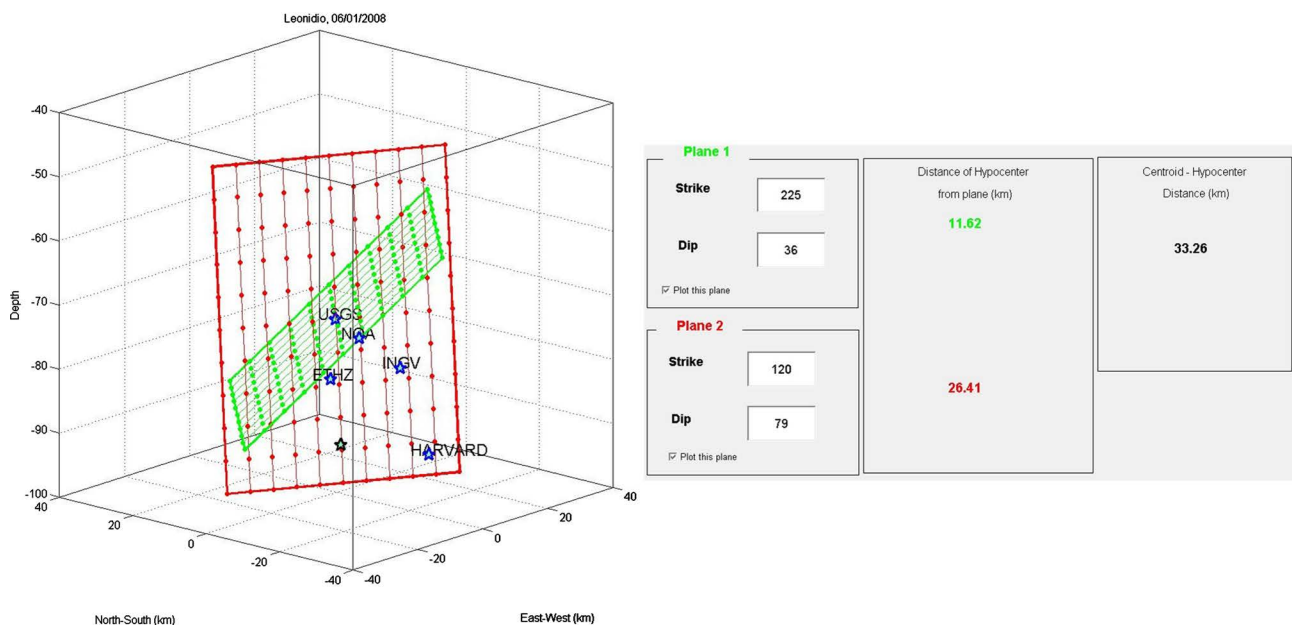
The fault plane is the nodal plane with strike = 225°, dip = 36° and strike = 19°. The distance of the hypocenter from this plane is 11.62 km, while the distance from the other plane is 26.30 km. The distance between Hypocenter and Centroid is 33.26 km (Figure 6). These results are in good agreement with this



**Figure 5.** Moment tensor solution of the 6 January 2008 (05:14 GMT) earthquake. The selected solution is highlighted with the green arrow in the misfit/CLVD-versus-depth diagrams. Down of this are appeared the summary of the solution and the corresponding beach ball. To the left of misfit/CLVD diagrams observed and synthetic displacement waveforms (continuous and dotted lines respectively) are shown, at the inverted stations for the radial, tangential and vertical components. At the lower part of the figure the summary of the solution and the fault plane solution as lower hemisphere equal-area projection, are depicted.

**Table 2.** Manually locations and moment tensor solutions by various agencies. Sources: HARV, INGV, USGS and ETHZ solution.

Agency	Lat (°)	Lon (°)	Depth (km)	M <sub>w</sub>	Strike 1 (°)	Dip 1 (°)	Rake 1 (°)	Strike 2 (°)	Dip 2 (°)	Rake 2 (°)
This study	37.11	22.75	86	6.0	120	79	125	225	36	19
HARV	36.98	22.87	92	6.1	117	77	130	222	41	20
USGS	37.30	23.00	73	6.2	106	79	127	210	38	18
ETHZ	37.10	22.70	70	6.2	108	74	116	228	30	34
INGV	37.00	22.80	73	6.2	113	76	131	219	43	21
UOA	37.10	22.66	75	6.0	114	75	120	228	33	28



**Figure 6.** Nodal planes 1 and 2 are shown in green and red rectangular, respectively. The hypocenter solutions of HARV, USGS, INGV, ETHZ and NOAA are shown in blue triangle.

one obtained by Zahradnik [2] [3].

### 3.3. The Mw = 6.7, and 6.1, 2008 Methoni Earthquakes

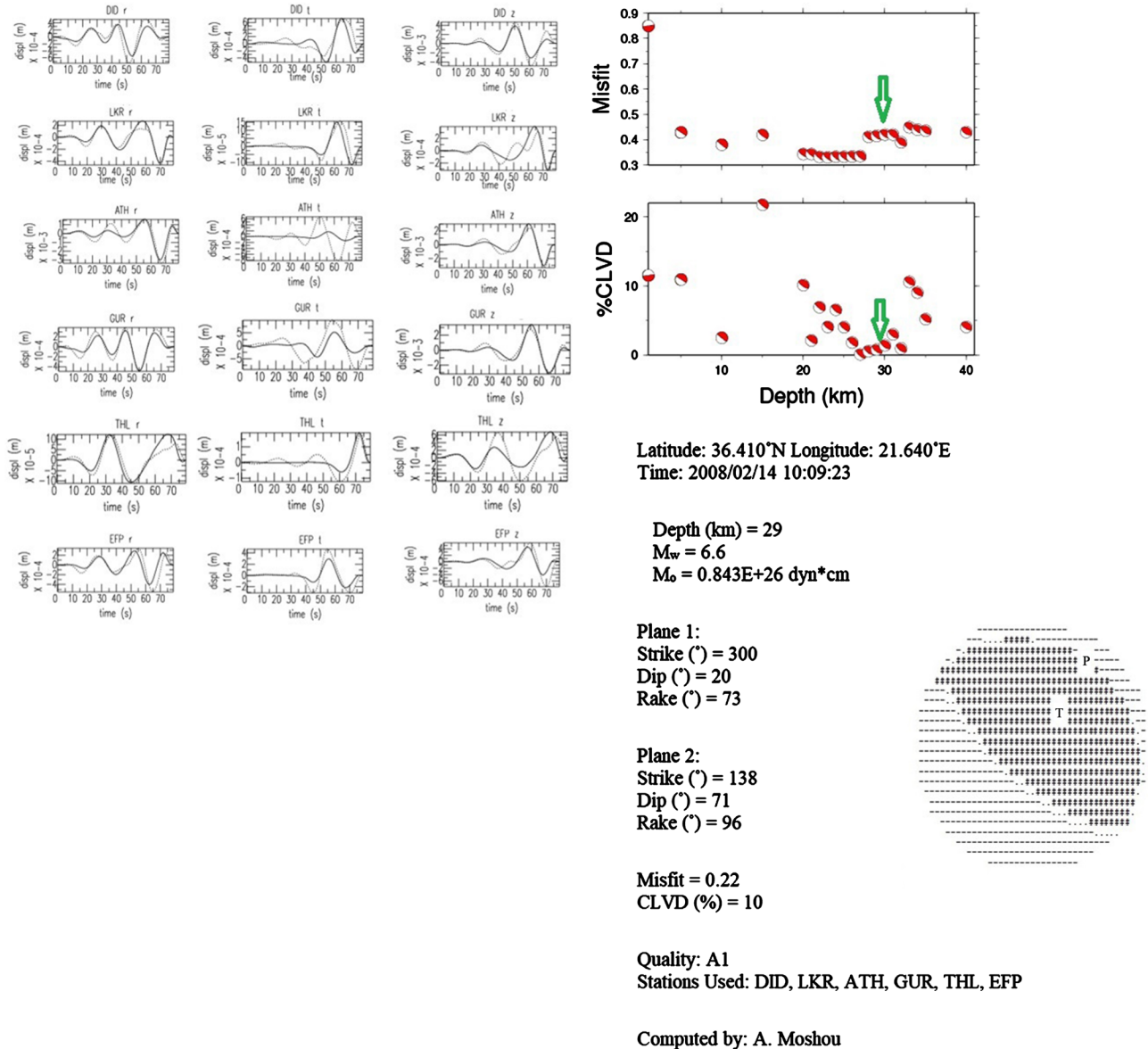
On February 2008 an earthquake sequence including three strong events (M<sub>w</sub> = 6.7, 6.1 and 6.0) occurred South of Methoni town, at a segment of the Hellenic arc which was not activated during the instrumental period. This sequence was followed by a large number of aftershocks, the strongest of which were processed to calculate their source parameters.

The first one occurred on 14 February 2008 (10:09, UTC) and the epicenter was located (36.50°N, 21.78°E) 230 km southwest of Athens. Two hours later (12:08, UTC) the second one occurred close to the first with epicenter (36.22°N, 21.75°E) and magnitude M<sub>w</sub> = 6.1, according to the National Observatory of Athens. These two strong events were followed by another earthquake six days later on 20 February 2014 with a similar magnitude (M<sub>w</sub> = 6.0). The moment tensor inversion indicates for the two first earthquakes the activation of a thrust-



faulting type, while for the third event the inversion indicates a strike-slip type faulting.

Following we present the result of the inversion for the first event. For this purpose, the data of 6 stations from Hellenic Unified Seismological Network (HUSN) at epicentral distances less than 360 km, were used. Reverse type faulting was revealed after applying inversion. The obtained focal mechanism is strike = 290°, dip = 16° and rake = 69°. The seismic moment is equal to  $M_0 = 1.56 \times 10^{26}$  dyn-cm, for a focal depth equal to 29 km (Figure 7). The inversion

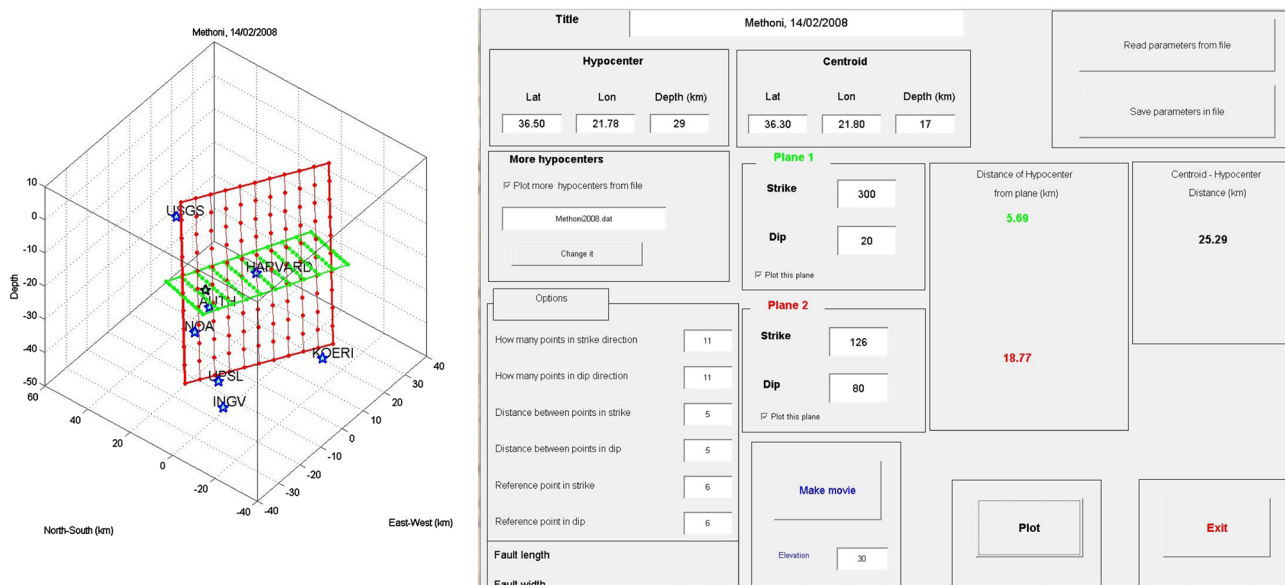


**Figure 7.** Moment tensor solution of the 14 February 2008 (10:09, UTC) earthquake. The selected solution is highlighted with the green arrow in the misfit/CLVD-versus-depth diagrams. Down of this are appeared the summary of the solution and the corresponding beach ball. To the left of misfit/CLVD diagrams observed and synthetic displacement waveforms (continuous and dotted lines respectively) are shown, at the inverted stations for the radial, tangential and vertical components. At the lower part of the figure the summary of the solution and the fault plane solution as lower hemisphere equal-area projection, is depicted.

resulted in a DC equal to 90%, while the compensated linear vector dipole was equal to 10%. To determine the main fault four representative solutions selected, are presented in **Table 3** and are plotted later in **Figure 8**.

**Table 3.** Manually locations and Moment Tensor Solutions by various agencies. Sources: HARV, INGV, USGS and AUTH solution.

Agency	Lat (°)	Lon (°)	Depth (km)	M <sub>w</sub>	Strike 1 (°)	Dip 1 (°)	Rake 1 (°)	Strike 2 (°)	Dip 2 (°)	Rake 2 (°)
This study	36.50	21.78	29	300	20	73	138	71	96	36.50
HARV	36.30	21.80	17	331	6	117	124	85	87	36.30
USGS	36.80	22.00	29	126	80	87	323	10	107	36.80
INGV	36.20	21.50	39	333	17	111	131	74	84	36.20
AUTH	36.50	21.80	35	131	77	92	302	13	81	36.50
UOA	36.30	21.69	35	290	16	69	132	75	96	36.30



**Figure 8.** H-C Plot using the UPSL hypocenter (green star), hypocenters of other agencies (USGS, HRV, AUTH, NOA, UPSL and INGV, blue stars).

### 3.4. The M<sub>w</sub> = 6.4, 2008 Andravida Earthquake

On June 8, 2008 (12:25 GMT) a strong earthquake with a moment magnitude M<sub>w</sub> = 6.4 occurred in the region of Peloponnese, W. Greece close to Andravida city. The area is located between the Gulf of Corinth extensional province and the Eurasia-Africa plate boundary offshore Cephalonia and Zakynthos Islands. No surface rupture was observed.

The epicentral coordinates were 37.98°N, 21.51°E, after these events a large number of aftershocks were followed. Double-difference relocations of 370 aftershocks show a linear pattern of events and define a clear NE-SW striking mainshock fault plane (Ganas *et al.*, 2009). The hypocenter was determined at 18 km depth beneath village Mihoi in SW Achaia.

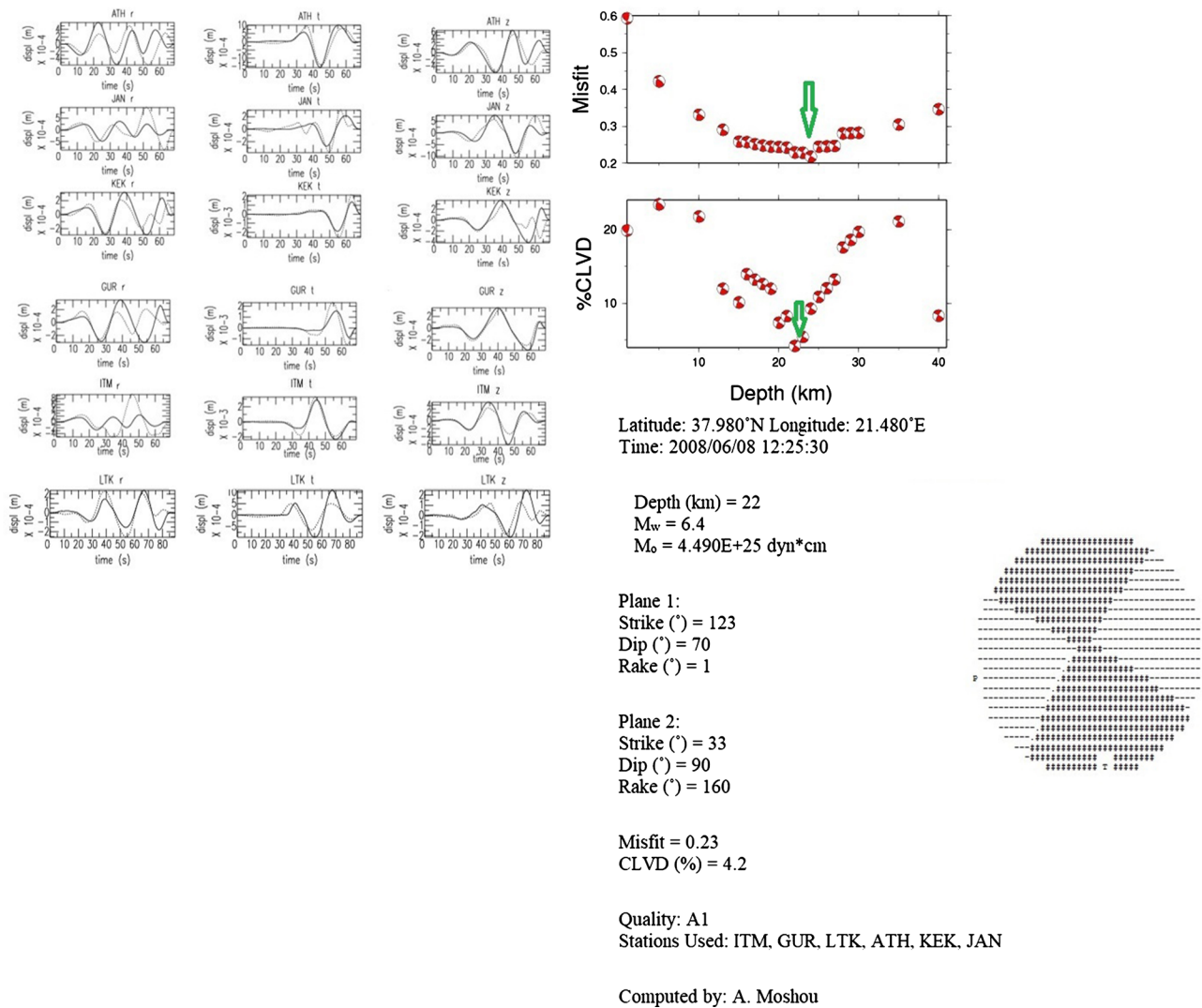
The constrained focal mechanism solution indicates strike-slip faulting with

source parameters strike = 290°, dip = 70° and rake = 1°, the seismic moment is equal to  $M_0 = 4.49 \times 10^{25}$  dyn-cm, while the focal depth was equal to 22 km (Figure 9). Taking to account the aftershock activity the fault plane is the one with the NE-SW direction.

To apply the HC-plot method, the following 6 hypocenters manually locations (Table 4) were considered.

The following CMT solutions were considered: HRV, USGS, ETHZ, INGV, AUTH and NOA (Figure 10). All these solutions are characterized by similar strike-dip-rake angles, with one nodal plane, strike~265°, dip~85° and rake~4° plotted in green and hereafter referred to as the green plane and the other one, strike~300°, dip~75° and rake~160° as the red plane (Figure 10).

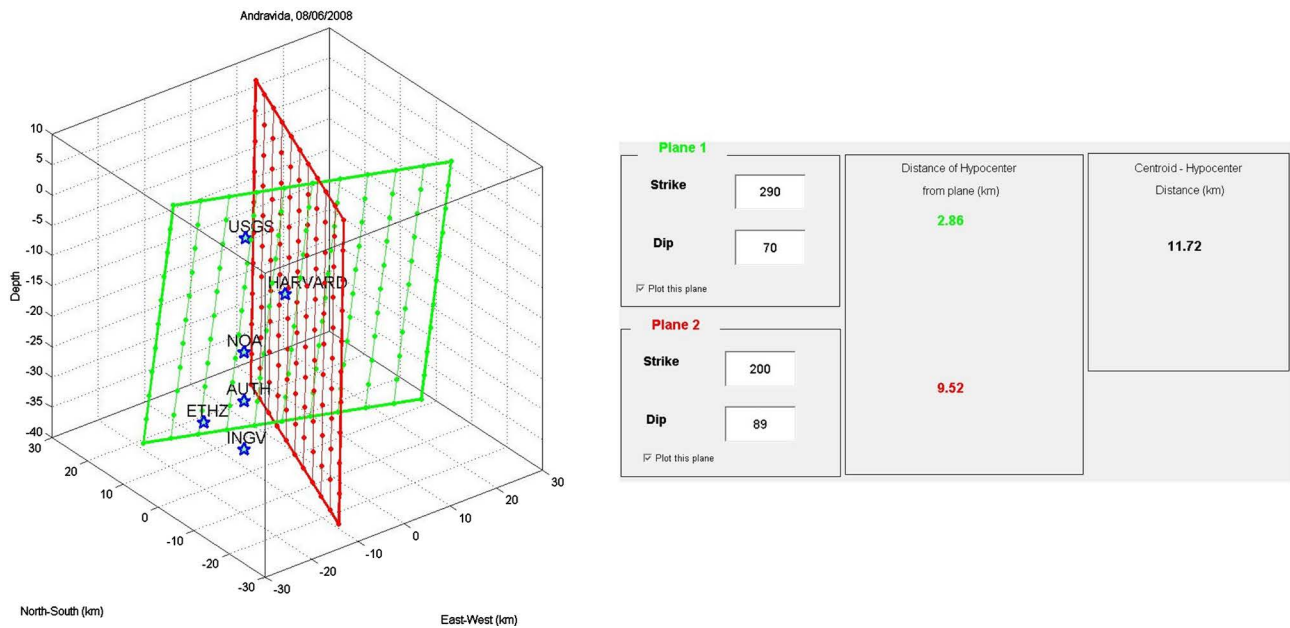
The fault plane is the nodal plane with strike = 290°, dip = 70° and strike = 1°.



**Figure 9.** Moment tensor solution of the event that occurred on 8 June 2008 (12:25, GMT). The selected solution is highlighted with the green arrow in the misfit/CLVD-versus-depth diagrams. Down of this are appeared the summary of the solution and the corresponding beach ball. To the left of misfit/CLVD diagrams we present observed and synthetic displacement waveforms (continuous and dotted lines respectively) at the inverted stations for the radial, tangential and vertical components.

**Table 4.** Hypocenter position from manually locations and focal mechanism by various agencies, by various agencies, <http://www.emsc-csem.org/Earthquake/tensors.php?id=88518&id2=j3155:INFO>.

Agency	Lat (°)	Lon (°)	Depth (km)	M <sub>w</sub>	Strike 1 (°)	Dip 1 (°)	Rake 1 (°)	Strike 2 (°)	Dip 2 (°)	Rake 2 (°)
This study	38.00	21.50	22	6.4	290	70	1	200	89	160
USGS	38.10	21.60	10	6.3	300	70	1	210	89	161
HARVARD	38.00	21.60	15	6.3	301	72	4	210	86	162
INGV	38.00	21.50	38	6.5	210	85	179	300	89	5
ETHZ	38.00	21.40	31	6.4	305	75	8	213	82	165
AUTH	38.00	21.50	30	6.5	211	90	178	301	88	1



**Figure 10.** Nodal planes 1 and 2 are shown in red and green, respectively. The depth that calculated by USGS, HARV, INGV, ETHZ and AUTH institutes are representing with blue triangle.

The distance of the hypocenter from this plane is 2.86 km, while the distance from the other plane is 9.52 km. The distance between Hypocenter and Centroid is 11.72 km (Figure 11). This result is in good agreement with those from other studies [2] [3] [13].

Following we present the inversions results for 28 events with magnitudes  $M_w > 3.5$  that occurred in the same region (Table 5).

### 3.5. The $M_w = 5.7$ , 2013 Limnos Earthquake

On 8 January 2013 at 14:16:08.3 UTC, a moderate earthquake of magnitude  $M_w = 5.7$  occurred off the southern coast of Limnos island. The event was strongly felt in nearby north Aegean islands, the neighboring Turkish coasts and the northeastern Greek mainland but caused no damage [14]. The epicenter was manually located at 39.6663°N, 25.5620°E, depth = 31 km with a local magnitude  $M_L = 5.8$  according to National Observatory of Athens,

[http://bbnet.gein.noa.gr/alerts\\_manual/2013/01/evman130108141608\\_info.html](http://bbnet.gein.noa.gr/alerts_manual/2013/01/evman130108141608_info.html).

The location of the earthquake indicates that it ruptured a fault segment running south of the North Aegean Trough near the island of Limnos [15] [16] [17], where the main, northern branch of the North Anatolian Fault (NAF) enters

**Table 5.** Moment Tensors Solutions for 28 studied events ( $M_w \geq 3.5$ ) of the Andravida Sequence (06/08/2008-18/06/2008). Columns (from left to right) show: event number, date and origin time, latitude and longitude of the epicenter, seismic moment, moment magnitude, the depth of the location and the depth as it calculated by the moment tensor inversion, strike dip and rake of the two nodal planes.

N/N	Origin Time		Location		Magnitude		$M_w$	Nodal Plane 1			Nodal Plane 2		
	Date (D/M/Y)	Time (UTC)	Lat (°)	Lon (°)	$M_0$ (dyn * cm)	Depth (km)		Strike (°)	Dip (°)	Rake (°)	Strike (°)	Dip (°)	Rake (°)
1	08/06/2008	12:25:28.180	37.9727	21.5157	4.60E+25	20.0	6.4	290.0	70.0	1.0	200	89	160
2	08/06/2008	12:40:53.958	38.1069	21.5651	4.30E+21	15.0	3.7	330.0	65.0	30.0	226	63	152
3	08/06/2008	12:43:39.173	38.0280	21.6107	1.60E+23	15.0	4.7	316.0	65.0	22.0	216	70	153
4	08/06/2008	12:46:53.112	38.0861	21.6454	7.00E+21	13.0	3.8	320.0	61.0	17.0	222	75	150
5	08/06/2008	12:55:13.978	38.1104	21.6325	2.80E+21	13.0	3.6	330.0	78.0	20.0	236	70	167
6	08/06/2008	13:20:30.038	38.0551	21.5630	2.10E+21	10.0	3.6	290.0	72.0	35.0	188	57	158
7	08/06/2008	13:27:49.114	38.0958	21.5940	2.80E+21	18.0	3.6	320.0	61.0	17.0	222	75	150
8	08/06/2008	13:31:34.668	38.0129	21.5616	7.60E+21	25.0	3.9	356.0	65.0	27.0	254	66	152
9	08/06/2008	13:41:47.828	37.9734	21.5114	6.80E+21	43.0	3.8	275.0	35.0	45.0	146	66	116
10	08/06/2008	14:01:07.601	38.0868	21.5968	6.81E+21	20.0	3.8	300.0	75.0	-15.0	34	76	-164
11	08/06/2008	14:16:25.262	38.0975	21.5783	3.70E+21	21.0	3.6	300.0	55.0	30.0	192	66	141
12	08/06/2008	14:22:29.173	38.0676	21.5783	5.50E+21	20.0	3.8	300.0	75.0	-15.0	34	76	-164
13	08/06/2008	16:10:06.321	37.9951	21.5388	6.80E+21	15.0	3.8	30.0	90.0	150.0	120	60	1
14	08/06/2008	17:56:03.654	38.0857	21.5875	2.30E+21	15.0	3.5	330.0	60.0	29.0	225	65	147
15	08/06/2008	18:12:39.103	37.9770	21.5234	2.70E+21	17.0	3.5	208.0	80.0	170.0	300	80	10
16	08/06/2008	18:42:25.763	38.0330	21.5549	3.76E+21	17.0	3.7	318.0	70.0	27.0	218	65	158
17	08/06/2008	20:36:23.571	38.1173	21.5926	2.40E+21	22.0	3.5	310.0	65.0	23.0	210	69	153
18	08/06/2008	21:10:22.743	38.0014	21.5788	3.90E+22	21.0	4.3	310.0	69.0	15.0	215	76	158
19	08/06/2008	21:48:30.423	37.9895	21.5296	7.93E+21	20.0	3.9	33.0	90.0	160.0	123	70	1
20	09/06/2008	01:32:03.476	38.0639	21.5667	3.70E+21	21.0	4.2	330.0	55.0	30.0	220	67	136
21	09/06/2008	11:51:59.347	38.1536	21.5830	2.80E+21	19.0	3.5	315.0	72.0	19.0	219	72	161
22	09/06/2008	13:53:21.609	38.0312	21.5513	2.30E+22	18.0	4.2	316.0	65.0	22.0	216	70	153
23	09/06/2008	16:18:40.590	38.0713	21.5678	1.90E+22	18.0	4.1	320.0	72.0	22.0	223	69	161
24	10/06/2008	20:45:17.632	38.0444	21.5626	3.76E+21	18.0	3.7	300.0	55.0	30.0	192	66	141
25	11/06/2008	01:05:03.649	37.9505	21.4908	2.80E+21	22.0	3.5	320.0	61.0	17.0	222	75	150
26	11/06/2008	16:09:43.337	38.0147	21.5562	2.80E+21	17.0	3.5	316.0	65.0	22.0	216	70	153
27	12/06/2008	03:15:33.988	38.0132	21.6093	2.80E+21	22.0	4.2	326.0	70.0	20.0	229	71	159
28	18/06/2008	17:57:26.520	38.1053	21.5921	4.30E+21	15.0	3.7	330.0	60.0	28.0	225	66	147

Aegean Sea [18] [19] [20].

The main event of January 8, 2013, was followed by a large number of aftershocks. For the next two months, a total number of 495 events ( $1 \leq M_L \leq 4.5$ ) were recorded and analyzed. The main shock, as well as the aftershocks were re-located [21]. The aftershock distribution of the mainshock reveals a NE-SW striking fault about 40 km offshore Limnos Island that extends from 2 km up to a depth of 14 km [21]. The next two months 7 events with magnitude  $M_w \geq 3.7$  occurred and the source parameters of them calculated and presented in **Table 6**.

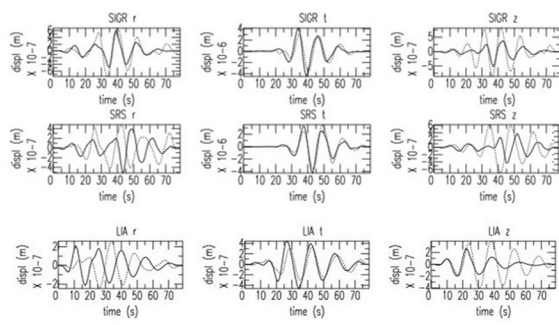
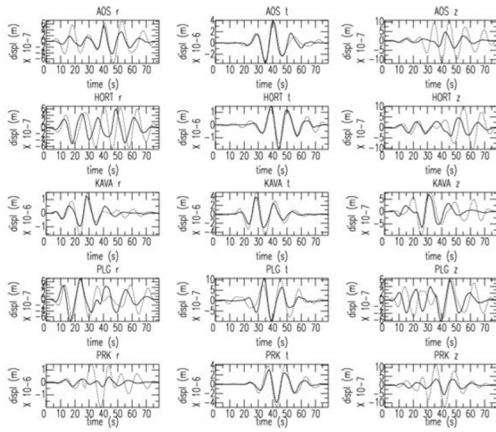
We applied moment tensor inversion to calculate the source parameters of the main event (January 8, 2013 14:16, UTC). Seismological data from HUSN, of 8 stations at epicentral distances less than 350 km were used. A band-passed filtering at frequencies 0.05 - 0.08 Hz was used both of recorded waveforms of three components and calculated synthetics seismograms. The inversion indicated a strike slip faulting, and the source parameters were calculated: strike =  $315^\circ$ , dip =  $86^\circ$ , rake =  $5^\circ$  with a depth 8 km while the moment magnitude determined  $M_0 = 3.90 \times 10^{24}$  dyn-cm. The calculated double couple found 95% (**Figure 11**).

To apply the HC-plot method, the followed 7 hypocenters manually locations (**Table 7**) were considered.

The following CMT solutions were considered: HRV, USGS, GFZ, INGV, AUTH and UOA (**Figure 13**). All these solutions are characterized by similar strike-dip-rake angles, with one nodal plane, strike $\sim 300^\circ$ , dip $\sim 80^\circ$  plotted in green and hereafter referred to as the “green” plane and the other one dipping, strike $\sim 65^\circ$  deg, dip $\sim 80^\circ$  as the “red” plane. The distance of hypocenter from this plane is 6.11 km, while the distance from the other plane is 1.61 km. The distance between Hypocenter and Centroid is 6.89 km (**Figure 12**). The main fault is this with source parameters: strike =  $45^\circ$ , dip =  $85^\circ$  and rake =  $-175^\circ$  and it is in agreement with these from the study [13].

**Table 6.** Moment Tensors Solutions for 7 studied events ( $M_w \geq 3.7$ ) of the Limnos Sequence (08/01/2013-08/03/2013), columns (from left to right) show: event number, date and origin time, latitude and longitude of the epicenter, seismic moment, moment magnitude, the depth as it calculated by the moment tensor inversion, strike dip and rake of the two nodal planes

NR	Origin Time		Location		Magnitude		Depth (km)	Nodal Plane 1			Nodal Plane 2		
	Date (D/M/Y)	Time (UTC)	Lat (°)	Lon (°)	$M_0$ (dyn*cm)	$M_w$	Inversion	Strike (°)	Dip (°)	Rake (°)	Strike (°)	Dip (°)	Rake (°)
1	09/01/2013	15:41:32.4	39.6872	25.5970	0.515E+23	4.4	7	88	63	-129	310	67	-25
2	10/01/2013	05:49:58.2	39.6618	25.5173	0.576E+22	3.8	6	68	56	-176	335	86	-34
3	11/01/2013	00:30:20.5	39.6623	25.5088	0.205E+23	4.2	9	331	78	-5	62	85	-168
4	11/01/2013	15:07:31.0	39.6982	25.5878	0.576E+22	3.8	8	60	89	-168	328	78	-4
5	13/01/2013	08:55:14.6	39.6627	25.5148	0.515E+23	4.4	7	58	71	-173	326	84	-19
6	13/01/2013	17:54:32.0	39.6462	25.6013	0.409E+22	3.7	7	60	75	-175	323	86	-23
7	05/03/2013	9:44:33.00	39.7023	25.5572	0.576E+22	3.8	8	64	50	-157	330	80	-30



Latitude: 39.550°N Longitude: 23.510°E  
 Time: 2013/01/08 14:16:08.3

Depth (km) = 8  
 $M_w = 5.7$   
 $M_0 = 2.90E+24 \text{ dyn}^2\text{cm}$

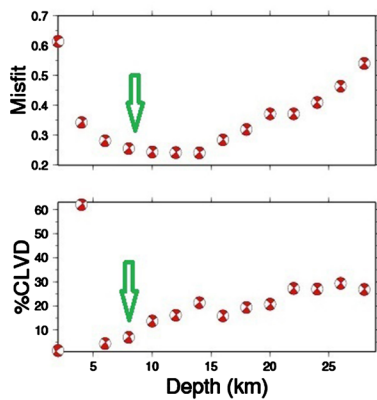
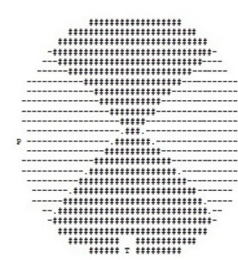
Plane 1:  
 Strike (°) = 315  
 Dip (°) = 86  
 Rake (°) = 5

Plane 2:  
 Strike (°) = 224  
 Dip (°) = 84  
 Rake (°) = 176

Misfit = 0.22  
 CLVD (%) = 5.0

Quality: A1  
 Stations Used: AOS, HORT, KAVA, PLG, PRK, SGR, SRS, LIA

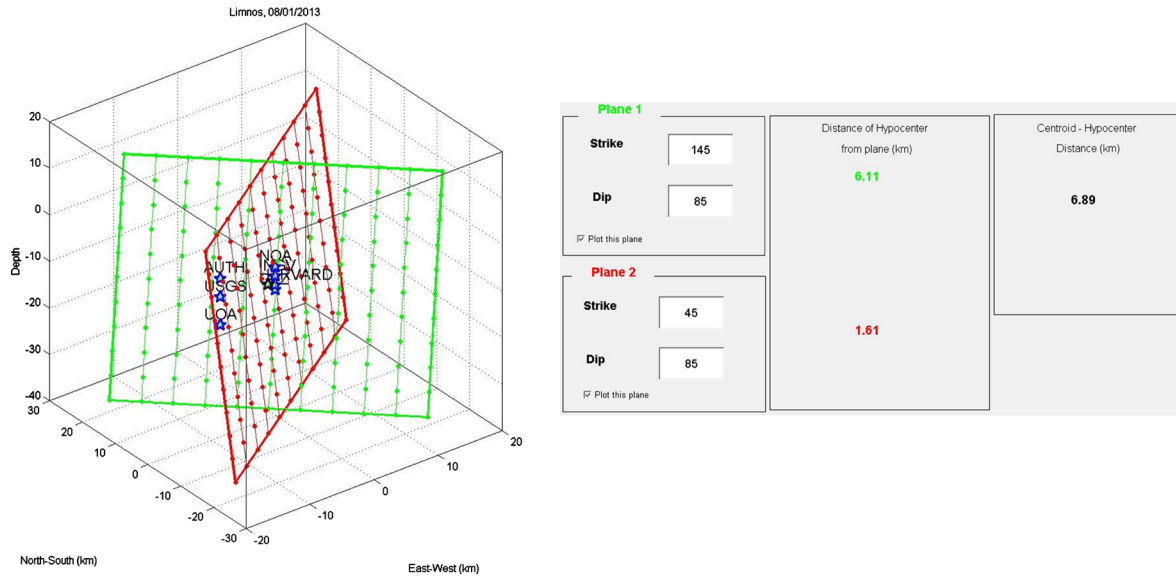
Computed by: A. Moshou



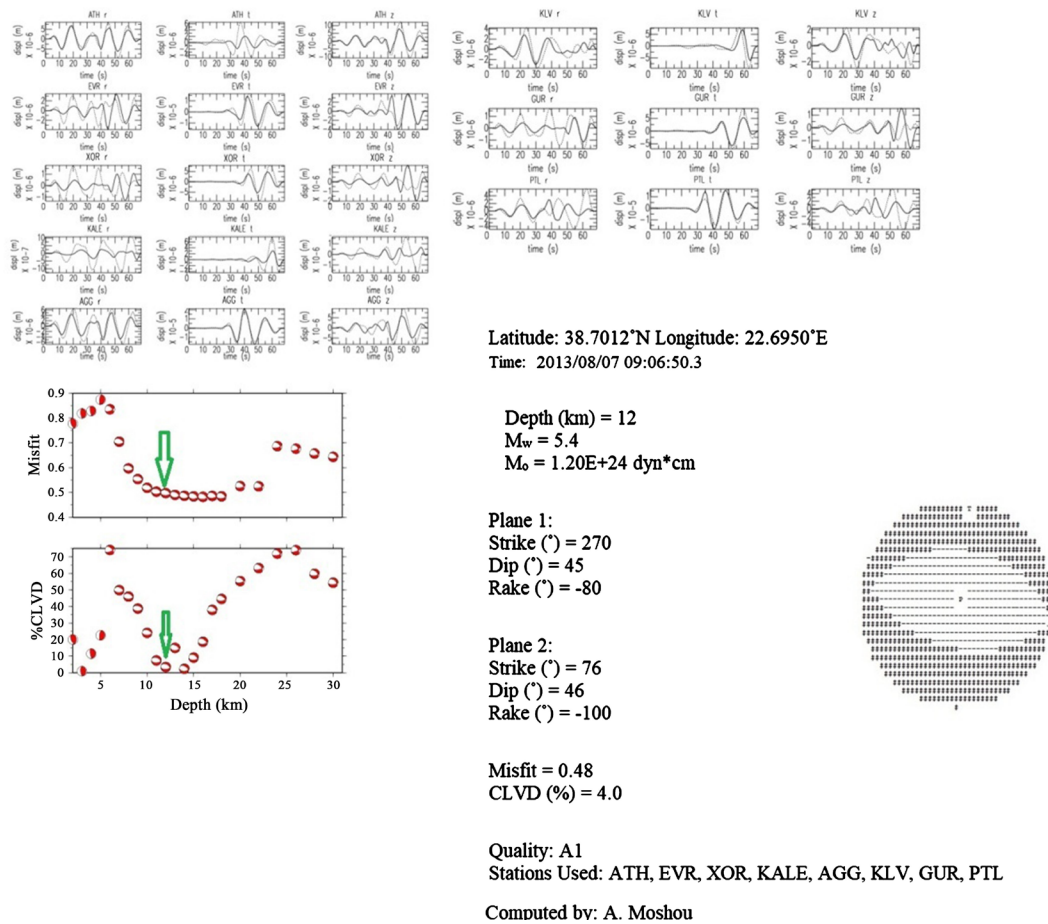
**Figure 11.** Moment tensor solution of the event that occurred on 8 January 2013 (14:16, UTC). Observed and synthetic displacement waveforms (continuous and dotted lines respectively) at the inverted stations for the radial, tangential and vertical components are appeared at the up section of the figure. Down of these, are presented (from left to right) the selected solution (highlighted with the green arrow) in the misfit/CLVD-versus-depth diagrams and the summary of the solution and the corresponding beach ball.

**Table 7.** Hypocenter position (manually locations) and Focal Mechanisms by various agencies <http://www.emsc-csem.org/Earthquake/tensors.php?id=300037&id2=MVT29;INGF>.

Agency	Lat (°)	Lon (°)	Depth (km)	Strike 1 (°)	Dip 1 (°)	Rake 1 (°)	Strike 2 (°)	Dip 2 (°)	Rake 2 (°)
This study	39.66	25.56	8	315	86	-5	45	85	-175
USGS	39.70	25.50	10	330	85	-12	61	78	-175
HARVARD	39.70	25.60	12	239	80	178	329	88	10
INGV	39.70	25.60	10	330	59	-9	65	82	-149
GFZ	39.70	25.60	13	331	77	-1	61	89	-167
UOA	39.70	25.50	16	331	78	-5	62	85	-168
AUTH	39.70	25.50	6	329	84	-1	59	89	-174



**Figure 12.** H-C Plot using the USGS, HARV, ING, GFZ, UOA and AUTH epicenters (with varying depth), and combining it with the CMT solution of HRV, for the earthquake of January 8, 2013 Limnos island.



**Figure 13.** Moment tensor solution of the event that occurred on 7 August 2013 (09:30, UTC). Observed and synthetic displacement waveforms (continuous and dotted lines respectively) at the inverted stations for the radial, tangential and vertical components. Down of these, are presented (from left to right) the selected solution is highlighted with the green arrow in the misfit/CLVD-versus-depth diagrams and the summary of the solution and the corresponding beach ball.



### 3.6. The $M_w = 5.4$ , 2013 Kalidromo Mountain Earthquake, Central Greece

On August 7, 2013 (09:06, UTC) a moderate earthquake,  $M_w = 5.4$  occurred in central Kallidromon Mountain, in the Pthiotis region of central Greece. The event was manually located at  $38.7012^\circ$  N,  $22.6950^\circ$  E, and the depth calculated at 8 km, using recordings of the HUS Network. Four minutes before at 09:02, (UTC) a foreshock of  $M_w = 4.3$  has occurred. For the next three months, eleven aftershocks with  $M_L = 4.0$  happened in the same area with the largest of these on September 16, 2013 (15:01, UTC) with  $M_w = 5.3$ . For the same time interval, 2270 aftershocks ( $1 < M_L < 5.2$ ) were recorded and relocated [22]. The moment tensor solutions for events with magnitudes,  $M_w \geq 3.7$ , were calculated and presented in the following table (Table 8).

**Table 8.** Moment Tensors Solutions for 24 studied events ( $M_w \geq 3.7$ ) of the Kalidromo Sequence (07/08/2013-07/11/2013). Columns (from left to right) show: event number, date and origin time, latitude and longitude of the epicenter, seismic moment, moment magnitude, the depth of the location and the depth as it calculated by the moment tensor inversion, strike dip and rake of the two nodal planes.

N/N	Origin Time		Location		Magnitude		$M_w$	Nodal Plane 1			Nodal Plane 2		
	Date (D/M/Y)	Time (UTC)	Lat (°)	Lon (°)	$M_0$ (dyn * cm)	Depth (km)		Strike (°)	Dip (°)	Rake (°)	Strike (°)	Dip (°)	Rake (°)
1	07/08/2013	09:02:45	38.7030	22.6676	2.76E+22	4.3	14.1	15	252	27	-87	69	63
2	07/08/2013	09:17:35	38.6857	22.6984	3.88E+21	3.7	11.3	9	275	35	-82	85	55
3	07/08/2013	09:18:13	38.6980	22.6926	3.88E+21	3.7	11.2	9	257	42	-86	72	48
4	07/08/2013	09:51:51	38.6863	22.662	3.88E+21	3.7	12.9	7	265	45	-90	85	45
5	07/08/2013	09:56:35	38.7013	22.695	1.84E+21	4.1	13.6	10	278	41	-86	93	49
6	07/08/2013	10:02:33	38.6817	22.685	3.88E+21	3.7	18.1	8	265	30	-87	82	60
7	07/08/2013	13:44:32	38.6908	22.690	1.15E+23	4.7	15.0	8	255	25	-80	64	65
8	07/08/2013	14:36:11	38.6862	22.683	3.88E+21	3.7	11.4	11	267	27	-97	95	63
9	07/08/2013	16:03:44	38.6863	22.647	3.88E+21	3.7	10.5	7	271	45	-89	90	45
10	08/08/2013	16:35:31	38.7102	22.683	3.88E+21	3.7	11.3	10	258	41	-88	75	49
11	09/08/2013	11:49:23	38.7010	22.772	1.50E+23	4.8	18.6	8	260	55	-100	97	36
12	09/08/2013	11:49:56	38.6937	22.677	1.00E+23	4.5	9.40	8	262	30	-86	77	60
13	09/08/2013	13:10:10	38.6915	22.647	1.50E+23	4.7	16.6	13	300	30	-30	57	76
14	09/08/2013	13:43:44	38.6768	22.625	7.01E+21	3.8	10.3	9	275	32	-79	82	59
15	14/08/2013	17:12:57	38.6995	22.667	4.87E+21	3.8	19.5	8	245	30	-88	63	60
16	18/08/2013	10:42:54	38.7018	22.762	7.01E+21	3.8	13.5	10	250	45	-87	66	45
17	18/08/2013	16:39:21	38.7033	22.797	8.10E+21	3.9	13.2	9	262	36	-89	81	54
18	18/08/2013	22:16:19	38.7090	22.678	4.87E+21	3.8	10.3	11	268	26	-96	95	64
19	16/09/2013	14:42:39	38.7000	22.768	1.04E+23	4.6	19.6	10	270	25	-93	93	65
20	16/09/2013	15:01:14	38.7188	22.753	8.66E+23	5.3	17.4	7	272	30	-90	92	60
21	16/09/2013	15:15:46	38.7118	22.790	3.88E+21	3.7	12.5	8	275	29	-86	90	61
22	16/09/2013	16:13:09	38.7082	22.718	2.94E+21	3.6	12.4	9	265	35	-87	81	55
23	17/09/2013	05:46:52	38.7210	22.730	7.01E+21	3.8	17.1	9	258	35	-85	72	55
24	17/09/2013	07:39:44	38.7015	22.777	1.84E+21	4.1	12.6	11	256	26	-105	93	65

Regional waveforms at epicentral distances less than  $3^\circ$  were used to determine the source parameters of the main event as well as all the aftershocks.

The inversion indicated the activation of a normal fault for all the studied events with a variation of the dip value between  $30^\circ$  and  $40^\circ$  and a focal depth varied between 8 - 13 km. For the first event (August, 7 2013 - 09:06 UTC) the regional data from 8 stations with good azimuthally coverage was used. The source parameters as they calculated applying the moment tensor inversion method found: strike =  $270^\circ$ , dip =  $45^\circ$ , rake =  $-80^\circ$  the seismic moment  $M_0 = 1.2 \times 10^{24}$  dyn-cm for a depth of 12 km. The calculated double couple was found equal to 96%, while the compensated linear vector dipole (CLVD) to 4%. The results of the applied procedure are presented in **Figure 13**.

To apply the HC-plot method, the followed 5 hypocenter manually locations (**Table 9**) were considered (**Figure 14**).

The previous solutions were considered: HRV, INGV, GFZ, AUTH and NOA. All these solutions are characterized by similar strike-dip-rake angles, with one nodal plane, strike $\sim 265^\circ$ , dip $\sim 50^\circ$ , plotted in green and hereafter referred to as the “green” plane and the other one dipping, strike $\sim 75^\circ$ , dip $\sim 50^\circ$ , as the “red” plane.

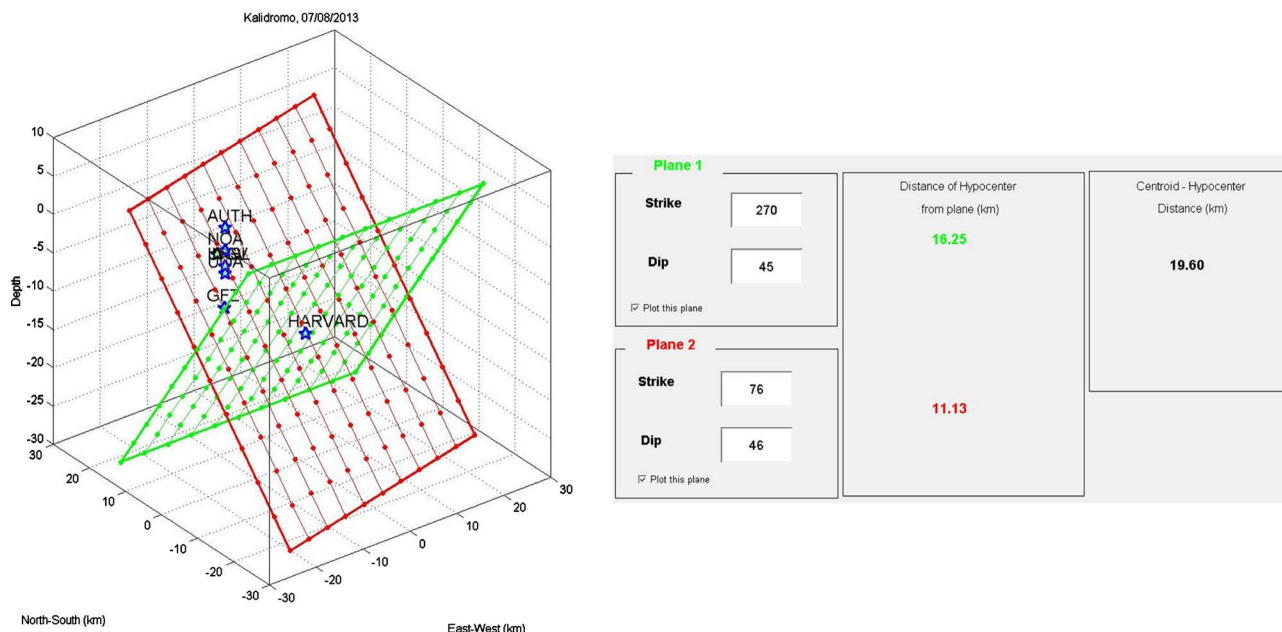
### 3.7. The $M_w = 6.1$ , 2014 Kefallinia Earthquake

On January 26, 2014 (13:55, UTC) two strong earthquakes of magnitude  $M_w = 6.1$  and  $M_w = 5.2$  (18:45, UTC) occurred on the island of Kefallinia, Ionia Sea. These events inducing extensive structural damages, mainly in the western and central parts. Eight days later on February 3, 2014 (03:08, UTC) a second strong event with a magnitude similar to the first ( $M_w = 6.0$ ) happened at the north section of Lixouri town. These two earthquakes ( $M_w = 6.1$  and  $M_w = 6.0$ ) occurred in the same island as the destructive events of 1953. Between August 9<sup>th</sup> and 12<sup>th</sup> of 1953, three earthquakes of magnitude 6.4, 6.8 and 7.2 took place in Cephalonia leading to hundreds of casualties and strong damages all over the island, but also in Zante and Ithaca.

These events were followed by a serious number of aftershocks. We noted that

**Table 9.** Hypocenter position (manually locations) and focal mechanisms by various agencies, <http://www.emsc-csem.org/Earthquake/tensors.php?id=328948&id2=OUE35:HARV>.

Agency	Lat ( $^\circ$ )	Lon ( $^\circ$ )	Depth (km)	Strike 1 ( $^\circ$ )	Dip 1 ( $^\circ$ )	Rake 1 ( $^\circ$ )	Strike 2 ( $^\circ$ )	Dip 2 ( $^\circ$ )	Rake 2 ( $^\circ$ )
This study	39.66	25.56	8	315	86	-5	45	85	-175
USGS	39.70	25.50	10	330	85	-12	61	78	-175
HARVARD	39.70	25.60	12	239	80	178	329	88	10
INGV	39.70	25.60	10	330	59	-9	65	82	-149
GFZ	39.70	25.60	13	331	77	-1	61	89	-167
UOA	39.70	25.50	16	331	78	-5	62	85	-168
AUTH	39.70	25.50	6	329	84	-1	59	89	-174



**Figure 14.** Nodal planes 1 and 2 are shown in red and green, respectively. Centroid is in the middle of the intersection of the nodal planes. The hypocenter solutions of USGS, HARVARD, INGV, GFZ, AUTH and NOAA are representing with blue triangle.

two months after the main event was recorded and analyzed more than 4000 events. The three large earthquakes, as well as all the aftershocks were relocated [1] [22] [23]. The source parameters for the strongest of these earthquakes ( $M_w \geq 4.0$ ) were calculated, using the moment tensor inversion and presented in **Table 10**.

Follow, we will present the focal mechanism for January 26, 2013 (13:55, UTC) earthquake,  $M_w = 6.1$  (**Figure 15**). The data of 6 stations at regional data in epicentral distances less than 370 km were used and inverted, as it appeared in **Figure 10**. The inversion indicated the activation of a strike-slip faulting with source parameters strike =  $23^\circ$ , dip =  $68^\circ$ , rake =  $175^\circ$  the depth calculated at 13 km and the moment magnitude  $M_0 = 1.51 \times 10^{25}$  dyn-cm. The fit between observed and synthetic waveforms and the misfit/CLVD versus depth diagram presented in **Figure 15**.

Six hypocenters manually locations were considered from NOAA, HARV, GFZ, INGV, UOA and AUTH (**Table 11**).

The following CMT solutions were considered: HRV, GFZ, INGV, AUTH, UOA and NOAA (**Figure 16**). All these solutions are characterized by similar strike-dip-rake angles, with one nodal plane, strike $\sim 20^\circ$ , dip $\sim 70^\circ$ , plotted in green and hereafter referred to as the “green” plane and the other one, with strike $\sim 130^\circ$ , dip $\sim 80^\circ$ , as the “red” plane.

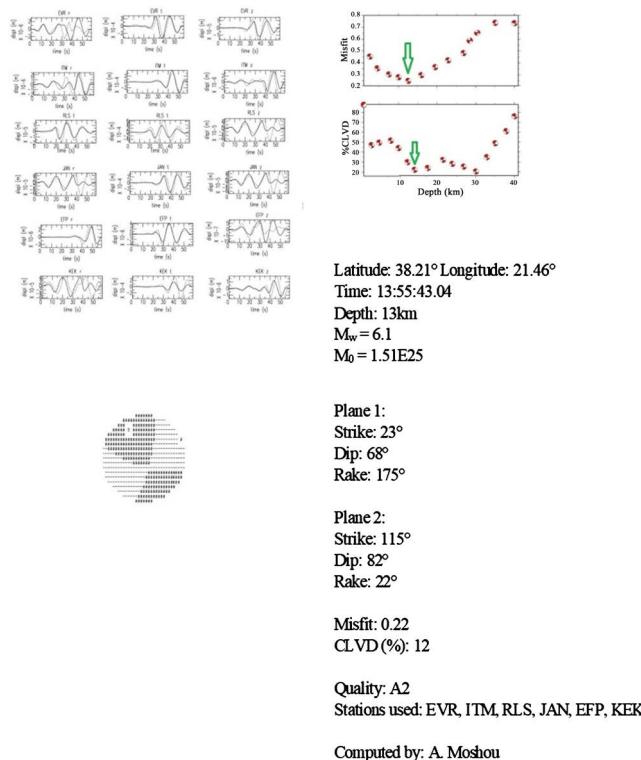
The distance of the hypocenter from this plane is 3.46 km, while the distance from the other plane is 3.87 km. The distance between Hypocenter and Centroid is 8.05 km (**Figure 17**). A preliminary result, about the activation of the main fault, is this with source parameters: strike =  $23^\circ$ , dip =  $68^\circ$  and rake =  $175^\circ$ . As it appears in **Figure 16** the distance between the plane and the Hypocenter-Centroid is similar.

**Table 10.** Moment Tensors Solutions for 36 studied events ( $M_w \geq 4.0$ ) of the Kefallinia Sequence (07/08/2013 - 05/03/2014). Columns (from left to right) show: event number, date and origin time, latitude and longitude of the epicenter, seismic moment, moment magnitude, the depth of the location and the depth as it calculated by the moment tensor inversion, strike dip and rake of the two nodal planes.

Nr	Origin		Location		Mo	Mw	Depth		Plane 1			Plane 2		
	Date (D/M/Y)	Time (UTC)	Lat (°)	Lon (°)	(dyn-cm)		Catalog	MT	Strike (°)	Dip (°)	Rake (°)	Strike (°)	Dip (°)	Rake (°)
1	26/01/2014	14:08:39	38.188	20.5325	2.96E+22	4.3	18.9	12	20	64	170	114	81	26
2	26/01/2014	14:21:58	38.2088	20.3787	9.52E+21	4.0	9.9	8	22	67	168	117	79	23
3	26/01/2014	14:24:04	38.2532	20.3903	2.22E+22	4.2	14.5	14	19	62	170	114	81	28
4	26/01/2014	14:41:39	38.2167	20.4757	2.22E+22	4.2	17.1	13	20	60	172	114	83	30
5	26/01/2014	14:55:50	38.2132	20.4100	9.52E+21	4.0	14.2	14	23	68	176	115	86	22
6	26/01/2014	14:59:25	38.3030	20.4753	7.35E+22	4.5	12.9	12	25	65	179	115	89	25
7	26/01/2014	15:36:39	38.2363	20.4373	1.38E+22	4.1	17.1	13	18	64	170	112	81	26
8	26/01/2014	19:03:07	38.1873	20.4177	2.96E+22	4.3	17.1	13	20	69	169	114	80	21
9	26/01/2014	19:12:04	38.2408	20.4002	4.06E+22	4.4	18.0	12	22	70	168	116	79	20
10	26/01/2014	21:15:34	38.1337	20.3002	8.17E+22	4.6	10.4	11	23	65	170	117	81	25
11	26/01/2014	21:42:12	38.1890	20.4862	9.52E+21	4.0	13.0	13	17	64	165	114	77	27
12	26/01/2014	23:06:55	38.2398	20.4297	2.22E+22	4.2	18.3	13	24	65	174	117	85	25
13	27/01/2014	09:47:38	38.1517	20.4025	1.38E+22	4.1	14.8	14	26	68	175	118	85	22
14	27/01/2014	13:05:50	38.2308	20.4403	2.96E+22	4.3	11.1	11	19	69	169	113	80	21
15	27/01/2014	15:39:34	38.3748	20.4222	2.22E+22	4.2	13.8	13	20	80	-173	289	83	-10
16	28/01/2014	01:05:55	38.2542	20.4347	9.52E+21	4.0	15.1	12	20	65	172	113	83	25
17	28/01/2014	05:12:53	38.2083	20.3817	2.96E+22	4.3	12.8	8	22	62	170	117	81	28
28	28/01/2014	08:07:11	38.2138	20.5502	9.52E+21	4.0	15.3	11	23	60	172	117	83	30
19	28/01/2014	14:49:33	38.2120	20.4552	9.52E+21	4.0	17.7	11	25	68	174	117	84	22
20	28/01/2014	19:12:11	38.4048	20.5022	1.38E+22	4.1	10.6	12	17	64	176	109	86	26
21	28/01/2014	22:22:37	38.4037	20.4885	2.22E+22	4.2	15.6	13	20	65	173	113	84	25
22	28/01/2014	22:23:39	38.3927	20.4418	2.22E+22	4.2	15.9	12	22	60	175	115	86	30
23	30/01/2014	11:06:18	38.4050	20.5267	4.06E+22	4.4	9.2	8	4	73	159	100	70	18
24	31/01/2014	06:52:47	38.4210	20.4843	4.06E+22	4.4	12.4	12	19	64	170	113	81	26
25	31/01/2014	12:45:40	38.4180	20.4677	2.96E+22	4.3	18.6	13	18	65	170	112	81	25
26	01/02/2014	16:33:38	38.1727	20.3876	7.35E+22	4.5	10.6	12	17	68	175	109	85	22
27	04/02/2014	19:42:12	38.2817	20.3702	2.22E+22	4.2	16.5	11	23	66	172	116	83	24
28	07/02/2014	03:26:43	38.3253	20.4325	2.22E+22	4.2	13	9	25	60	170	120	81	30
29	07/02/2014	08:59:43	38.2338	20.4558	2.22E+22	4.2	12.9	12	20	65	175	114	81	25
30	09/02/2014	08:22:58	38.1752	20.3675	7.35E+22	4.5	11.2	12	20	67	180	110	90	23

## Continued

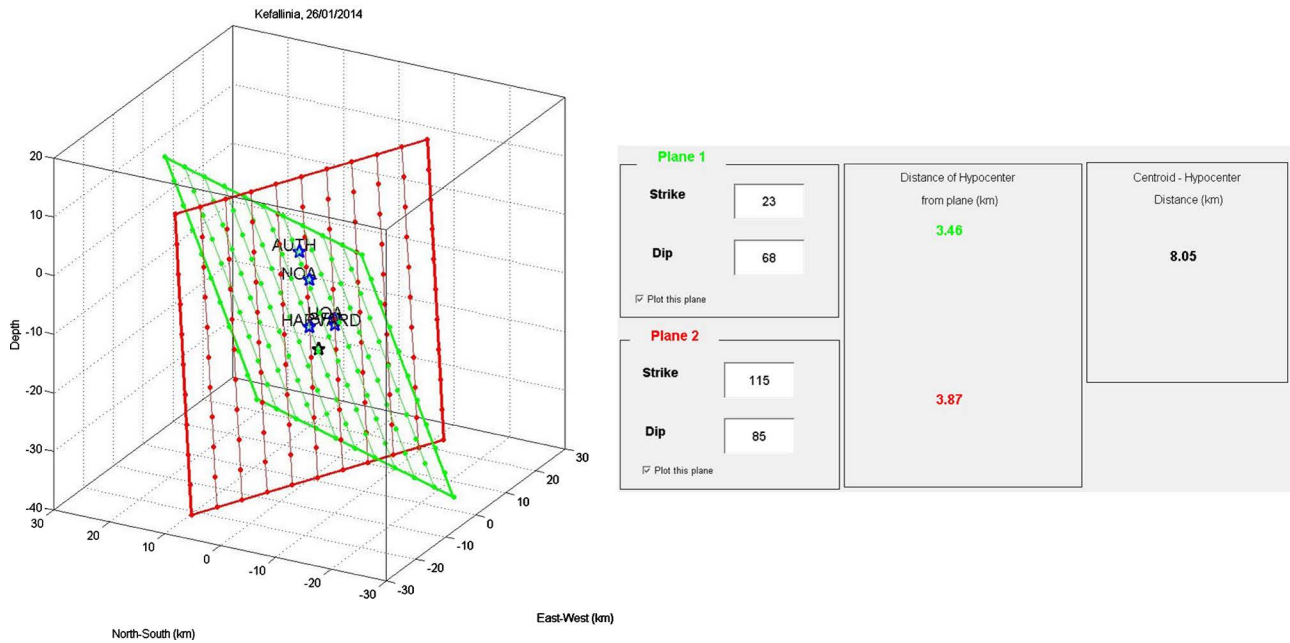
31	12/02/2014	10:34:31	38.1655	20.3538	8.17E+22	4.6	11.1	12	25	70	166	120	77	21
32	14/02/2014	03:38:33	38.1677	20.3432	8.17E+22	4.7	9.8	13	20	67	170	114	81	23
33	21/02/2014	15:18:23	38.2147	20.972	7.35E+22	4.5	16.2	13	26	73	177	117	87	17
34	05/03/2014	12:49:20	38.078	20.3092	8.17E+22	4.6	20.4	12	30	70	170	123	80	15
35	05/03/2014	15:08:43	38.0792	20.3467	1.38E+22	4.1	18	14	18	70	168	112	79	20
36	05/03/2014	18:42:02	38.1423	20.4185	9.52E+21	4.0	16.3	12	22	65	168	117	79	25



**Figure 15.** Moment tensor solution of the event that occurred on 26 January 2014 (13:55, UTC). Observed and synthetic displacement waveforms (continuous and dotted lines respectively) at the inverted stations for the radial, tangential and vertical components. Down of these, are presented (from left to right) the selected, solution is highlighted with the green arrow, in the misfit/CLVD-versus-depth diagrams and the summary of the solution and the corresponding beach ball.

**Table 11.** Hypocenter position (manually locations) and focal mechanisms by various agencies, <http://www.emsc-csem.org/Earthquake/tensors.php?id=357329&id2=kr068;INFO>.

Agency	Lat (°)	Lon (°)	Depth (km)	Strike 1 (°)	Dip 1 (°)	Rake 1 (°)	Strike 2 (°)	Dip 2 (°)	Rake 2 (°)
This study	38.21	21.46	18	23	68	175	115	85	22
HRV	38.20	20.40	12	12	45	154	120	72	48
GFZ	38.20	20.50	14	300	56	43	183	56	138
INGV	38.20	20.40	8	13	43	161	118	77	49
UOA	38.20	20.50	16	35	62	175	124	80	20
AUTH	38.20	20.50	7	17	87	-177	287	87	-3



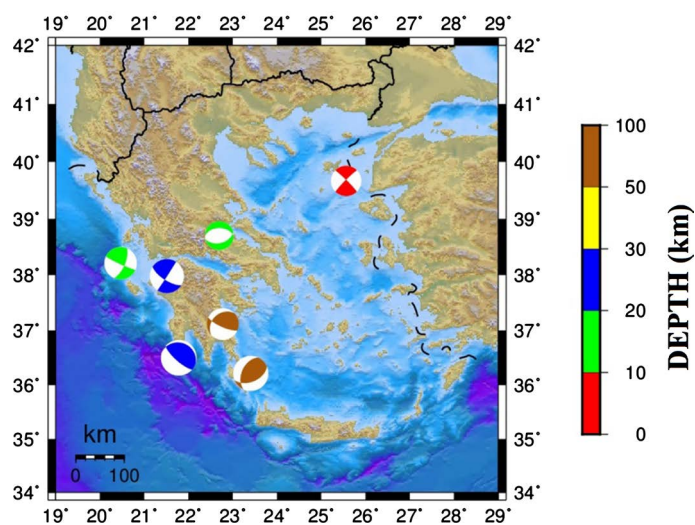
**Figure 16.** Nodal planes 1 and 2 are shown in red and green, respectively. The hypocenter solutions of USGS, HARVARD, INGV, ETHZ and AUTH are representing with blue triangle.

#### 4. Conclusions

The knowledge of the source parameters for moderate earthquakes is very important for seismically active regions, especially in the case where no large events occur. In general, it allows analytical studies; reveals the tectonics and the seismogenic characteristics of a specific region. The methodology that was used in this study is applied to the seven largest events that occurred recently in Greece. Concerning the 2006 Kythira and the 2008 Leonidio deep events at depths equal to 69 km and 85 km respectively [24] [25]. On the other hand, the seismic parameters of the 2008 Methoni and Andravida earthquakes, the 2013 Kalidromo Mountain (Central Greece) [26] and finally the most recent event the 2014 Kefallinia Island were determined using the proposed methodology. For this purpose, regional data from Hellenic Unified Seismological Network, with epicentral distances less than 350 km, were selected and analyzed. The methodology is based on the generation of synthetic seismograms using the method of Kennett [8] for given earth-structure and then they compared with the corresponding observed. The next step was the deconvolution of the instrument response from the waveforms and then their integration to produce displacement. A band-pass filter was applied both to synthetic and observed seismograms and finally horizontal components rotated to radial and transverse. For all the events with magnitude,  $M_w > 5.5$  that occurred in Greece at the last years the source parameters as well as the fault that re-activated was calculated and present in **Table 12**. With bold represent the main fault, after the text edit has been completed, the paper is ready for the template. Duplicate the template file by using the Save as command, and use the naming convention prescribed by your journal for the name

**Table 12.** Earthquakes source parameters determinate in the present study, <http://www.globalcmt.org/CMTsearch.html>.

N/N	Origin Time		Location		Magnitude		Depth (km)	Nodal Plane 1			Nodal Plane 2			Institute
	Date (D/M/Y)	Time (UTC)	Lat (°)	Lon (°)	M <sub>0</sub> (dyn*cm)	M <sub>w</sub>		Strike (°)	Dip (°)	Rake (°)	Strike (°)	Dip (°)	Rake (°)	
1	08/01/2006	11:34:54.00	36.214	23.406	8.40E+25	6.40	69	<b>205</b>	<b>48</b>	<b>59</b>	67	50	120	This study
		11:35:00.30	35.930	23.290	1.51E+26	6.70	64	201	44	55	66	55	119	HARVARD
2	06/01/2008	05:14:19.30	37.114	22.775	1.60E+25	6.00	85	114	75	120	<b>228</b>	<b>33</b>	<b>28</b>	This study
		05:14:23.70	36.980	22.870	2.25E+25	6.20	92	117	77	130	222	41	20	HARVARD
3	14/02/2008	10:09:23.40	36.500	21.780	8.43E+25	6.60	29	<b>290</b>	<b>16</b>	<b>69</b>	132	75	96	This study
		10:09:29.00	36.240	21.790	2.37E+26	6.80	20	332	6	120	121	85	87	HARVARD
4	08/06/2008	12:25:27.90	37.980	21.510	4.49E+25	6.40	22	<b>123</b>	<b>70</b>	<b>1</b>	33	90	160	This study
		12:25:36.90	37.930	21.630	4.56E+25	6.40	24	301	74	7	209	83	164	HARVARD
5	08/01/2013	14:16:08.32	39.663	25.562	3.90E+24	5.70	8	315	86	-5	<b>224</b>	<b>84</b>	<b>-176</b>	This study
		14:16:11.40	39.620	25.610	6.63E+24	5.80	15	331	83	-1	61	59	-173	HARVARD
6	07/08/2013	09:06:51.86	38.701	22.680	1.20E+24	5.40	12	270	45	-80	<b>76</b>	<b>46</b>	<b>-100</b>	This study
		09:06:54.00	38.540	22.690	1.60E+24	5.40	13	267	27	-78	73	64	-96	HARVARD
7	26/01/2014	13:55:43.04	38.219	20.532	1.51E+25	6.10	5	<b>23</b>	<b>68</b>	<b>175</b>	115	85	22	This study
		13:55:50.40	38.150	20.360	2.04E+25	6.10	14	20	65	177	111	87	25	HARVARD

**Figure 17.** Focal mechanisms solutions, determined using regional data, in this study between 2006-2014.

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The method was applied for a large variety of magnitudes and could determine the fault plane orientation and the seismic moment even under conditions of poor azimuthal coverage, as the fit of data and synthetics was well predicted for most events. All the solutions were compared with those from other institutes and they were in very good agreement. The focal mechanisms solutions for the events determined in this study appear in **Figure 17**.

## Acknowledgements

I acknowledge the use of Hellenic Unified Seismograph Network (HUSN) data and I would like to thank the NOA scientific personnel for phase picking. The open-source software GMT <http://www.soest.hawaii.edu/gmt/> was used to make several figures.

I gratefully thank the operators of the European permanent seismic networks who make their data available through EIDA, <http://www.orfeus-eu.org/eida>. In this study data from the following Institutes were used.

- HL (NOA, Hellenic Seismic Network), doi:10.7914/SN/HL.
- HT (Aristotle University of Thessaloniki Seismological Network), doi:10.7914/SN/HT.
- HP (University of Patras, Seismological Laboratory), doi:10.7914/SN/HP.
- HA (National and Kapodistrian University of Athens, Seismological Laboratory), doi:10.7914/SN/HA.

## Conflicts of Interest

The author declares no conflicts of interest regarding the publication of this paper.

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